

What is Albedo Susceptibility (S₀)



- S_o - Changes in cloud shortwave albedo (A_c) due to perturbations in cloud droplet number concentration (N_d).

- *Twomey effect* (timescale ~5-10mins) - added aerosol serves as extra cloud condensation nuclei, leading to more smaller droplets (more nuclei competing for same amount of condensate), assuming constant liquid water path (LWP). This leads to **brighter clouds.** - Subsequent changes in cloud condensate, e.g. LWP, can manifest ~20hrs after the perturbation. The **sign of S_0** depends on the LWP



Why Important to Understand and Quantify

- Routine ship emissions do not always lead to detectable **shiptracks**. - Both the Twomey effect and the LWP adjustment are dependent on the **cloud regimes** & the **large-scale meteorological conditions**. - It is important to characterize the cloud/meteorology regimes associated with the susceptible (and less susceptible) conditions, as well as their **geographical preference** & **frequency of occurrence**.





Goals

- Quantify albedo susceptibility of marine low clouds and produce a satellite-derived climatological susceptibility map.

- Distinguish **susceptibility regimes** (brightening & darkening) and quantify their **frequency of occurrence**.
- Associate susceptibility regimes with large-scale meteorological conditions: a focus on their **seasonal covariability**.



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Albedo Susceptibility Regimes in LWP-N_d Space



- +ve S₀ left of the $r_e = 12 \ \mu m$ isoline consistent with the precipitation-suppression induced **brightening** potential.

- -ve S_o right of evaporation-entrainment feedback (EEF) curve presents clouds with darkening potential.

- +ve S₀ for non-precip thin clouds (LWP<60 gm⁻²; weak entrainment velocity) presents the most frequently occurring **Twomey brightening** potential. [Twomey effect ~ (1-A_c)/3]







Seasonal Covariability between Albedo Susceptibility and Large-scale Meteorology

- Left column: Monthly mean radiative susceptibility (F₀; black) and frequency of occurrence of SLLC (red). Color of the circle indicates monthly mean LWP. Size of the circle indicates monthly mean N_d. Open (closed) circles indicate likely precipitating (non-precipitating) condition, based on calculated critical LWP values with $r_e = 12 \mu m$. - Middle column: Monthly mean F_o into 3 regimes: non-precipitating brightening (dotted blue), darkening (brown), and precipitating brightening (solid blue).

- Right column: Monthly mean meteorological conditions: LTS (blue), RH_{ft} (green), SST (red), and CTH (cyan).
- overall the most susceptible, especially Mar.-Sept. when LWP is low and N_d is relatively high. - What can we learn from seasonal covariabilities among meteorology, cloud frequency and albedo susceptibility
 - Pacific, SE Atlantic, Australian);
 - High susceptibilities can coincide with high N_d (e.g. SE Atlantic, NE Atlantic);
 - Distinctively covarying meteorological factors leads to drastically different susceptibility evolutions among global stratocumulus regions;

Take-home points

- entrainment feedback, and Twomey brightening processes. 60°N) marine low clouds, most pronounced over **subtropical coastal regions** where marine stratocumulus prevail.
- remote S.E. Pacific/Atlantic).

Distinct regional covariabilities between low-cloud albedo susceptibility and the large-scale meteorological conditions are observed over global marine stratocumulus regions.



- Geographical distributions of marine lowcloud (a) albedo susceptibility (S₀) and (b) the product of **radiative susceptibility (F_o)** and annual frequency of occurrence of single layer liquid cloud (SLLC), derived from the Aqua dataset (2005-2007).
- **Radiative Susceptibility (F**₀) is the product of S_o, f_c and downwelling SW flux; **+ve value** indicate cooling.
- An overall brightening potential over global oceans, more pronounced over subtropical coastal stratocumulus, especially when the frequency of occurrence of SLLC is taken into account (panel b).
- Geographical distributions of the **radiative** susceptibility (F_o) in 3 regimes: (a) nonprecipitating brightening, (b) darkening, and (c) precipitating brightening.
- The 3 regimes are separated based on the sign of S_o and a critical LWP value (calculated using an effective radius of 12 μ m) of all individual $1^{\circ} \times 1^{\circ}$ snapshots within a $5^{\circ} \times 5^{\circ}$ area. Only areas with SLLC frequency of occurrence greater than 0.1 are shown.
- Non-precip brightening dominates coastal regions, where annual-mean N_d is high and the MBL is shallow.
- Over the remote oceans where MBLs are deeper and clouds are thicker and mostly precipitating with lower f_c, **brightening** potential associated with rain suppression is the dominant regime.
- **Thicker marine stratocumulus** (high f_c and non-raining) that populate the remote parts of global subtropical stratocumulus decks show the strongest cloud darkening potential, especially over S.E. Pacific/Atlantic where it overcomes the brightening in annual mean.

- Among 5 subtropical stratocumulus regions, S.E. Pacific is the least susceptible, owing to a predominant cloud darkening regime that is consistent with the dry free-tropospheric, cool ocean surface, and the stable atmosphere; **N.E. Atlantic is**

- The most susceptible conditions/months can coincide with the least frequent warm cloud occurrence (e.g. NE Pacific, SE

- SST, LTS and RH_{ft} evolve **distinctively** even among major subsidence regions (i.e. SE Pacific, SE Atlantic and NE Pacific);

- Distinct **regimes** of albedo susceptibility appear in **LWP-N_d** variable space, consistent with **precip-suppression**,

- Satellite observations (CERES-MODIS) indicate an **overall cloud brightening** potential (annual mean) for global (60°S – - Entrainment driven **negative LWP adjustments** offset the Twomey & precip-suppression brightening over **remote oceans**

where deeper, thicker, mostly non-precipitating clouds populate (darkening **overcomes** brightening in annual mean over