



Investigating potential triggers for the Larsen B fast ice break-up event and the initial glacier response

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Introduction

In late March 2011, landfast sea ice started to form in the Larsen B embayment that persisted **until January 2022**. In the eleven years of continuous landfast ice presence, the Larsen B tributary glaciers developed extensive mélange areas and formed ice tongues that extended up to 10 km from the 2011 ice fronts. Breakout of the landfast ice began **20 January 2022**, leading immediately to retreat and break-up of the glacier mélange and ice tongue areas (Figure 1). The tributary glaciers have responded dynamically to changes in their stress regime in the past, such as the disintegration of the Larsen B Ice Shelf in 2002. Here we present our analysis of **potential triggers** for the loss of landfast ice in January 2022 and and an initial analysis of the glacier response to the event.

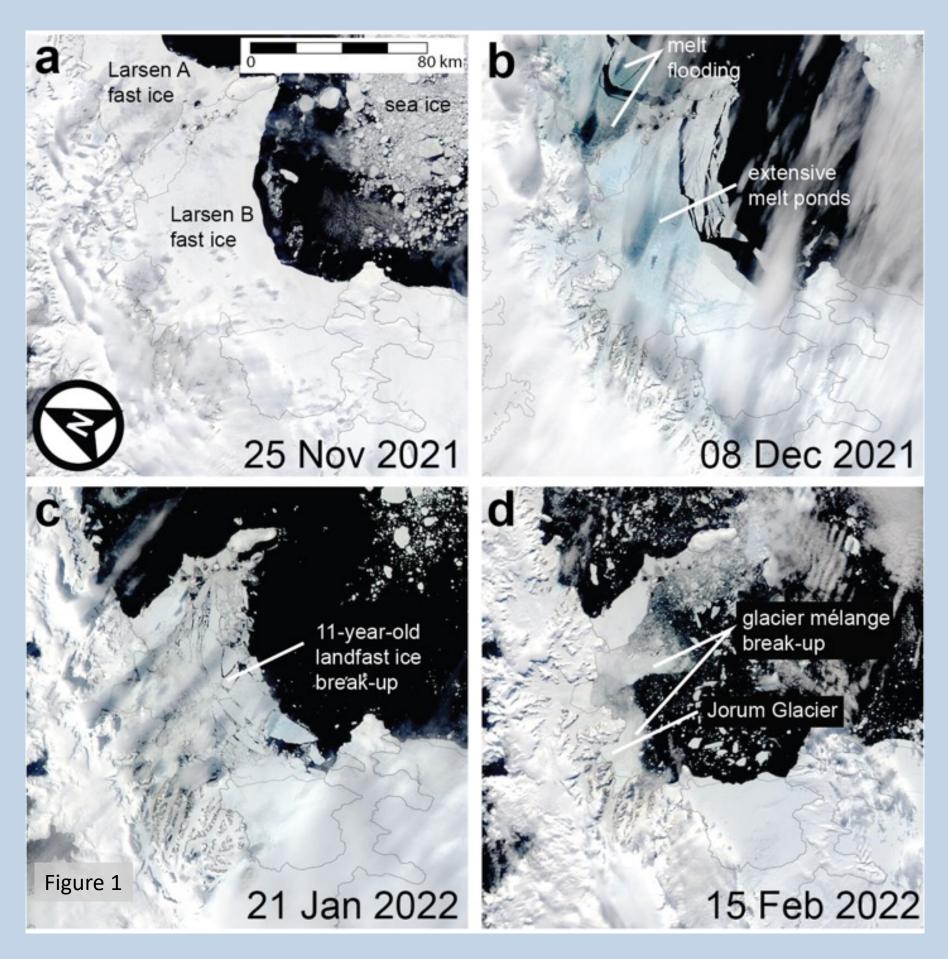
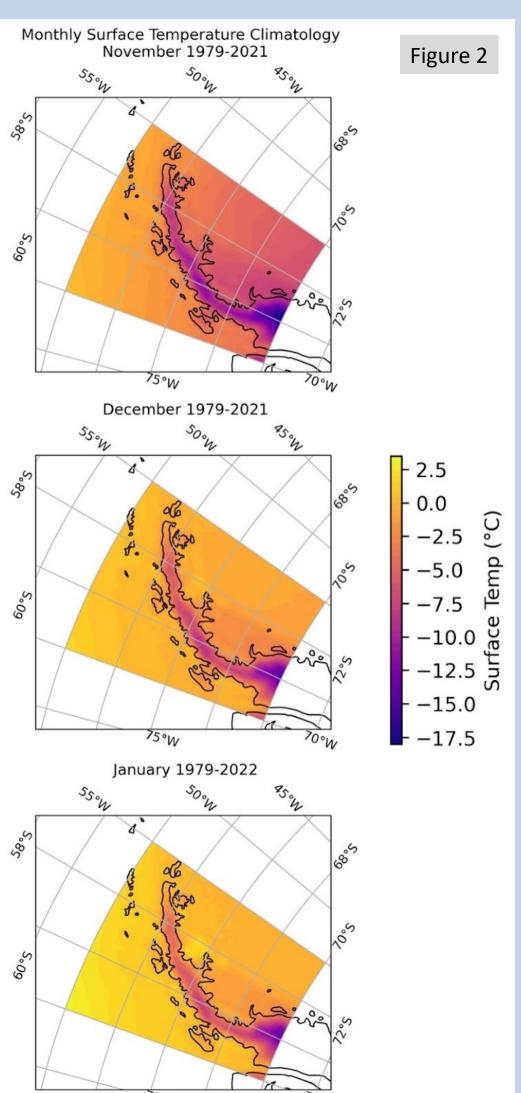


Figure 1: Landfast ice was present from the end of March 2011 until 20 January 2022 in the Larsen B Embayment. Extensive melt ponding and arcuate fracturing preceded this break-up event. Essentially overnight the landfast sea ice had fractured and flowed out of the embayment, exposing the tributary glaciers to the open ocean.



Setting the Stage: Climatology

Figure 2: NDJ spring-summer monthly climatology calculated from ERA-5 reanalysis data from 1979-2022 (43 years), using the averaged monthly 2m temperature product for months of interest.

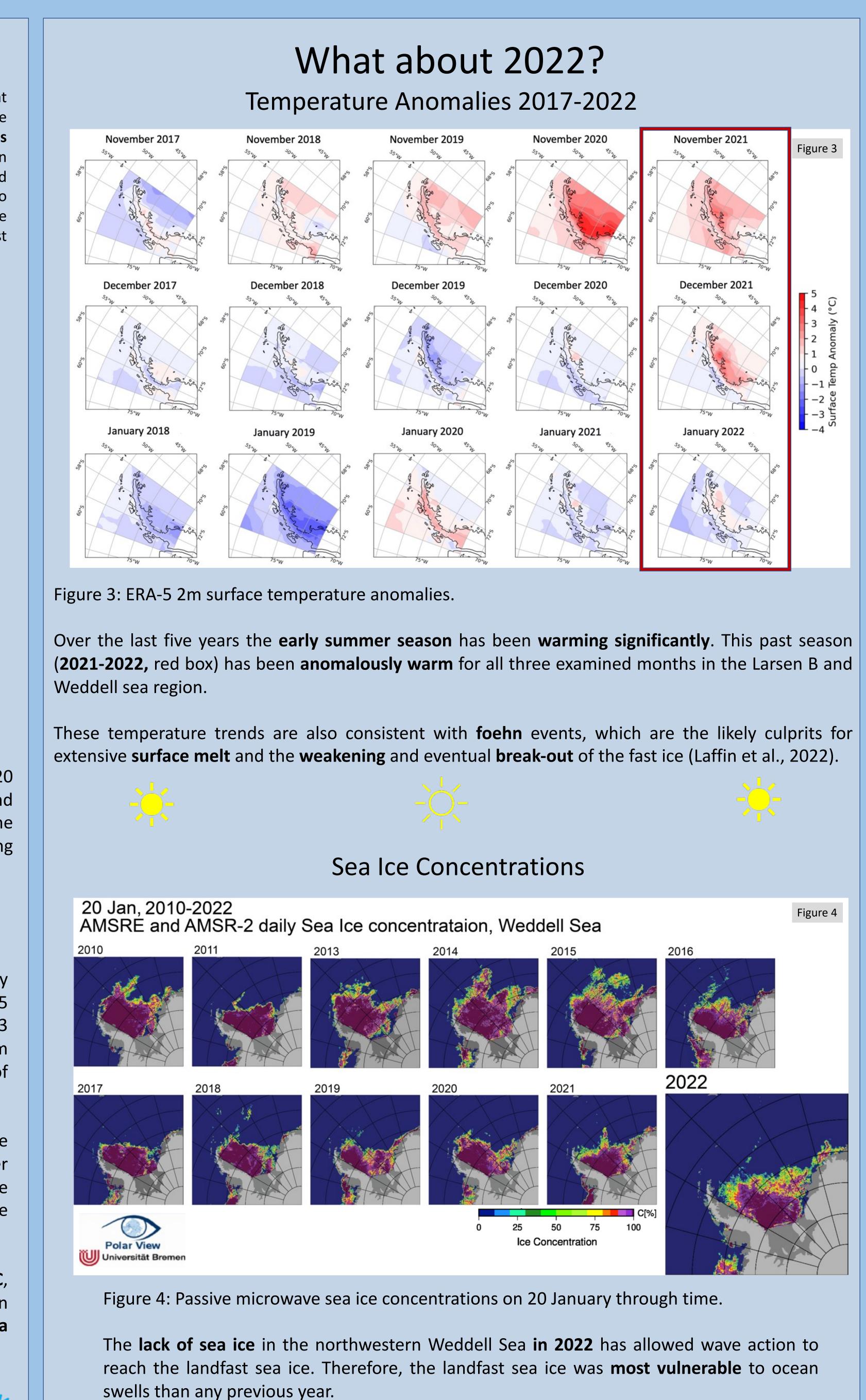
November is the **coolest** month while January is the warmest and December is between the two. Generally, the Amundsen Sea is warmer than the Weddell Sea.

Temperatures tend to be **below 0°C**, apart from the Larsen B and C region in January where the Weddell Sea exceeds 0°C.

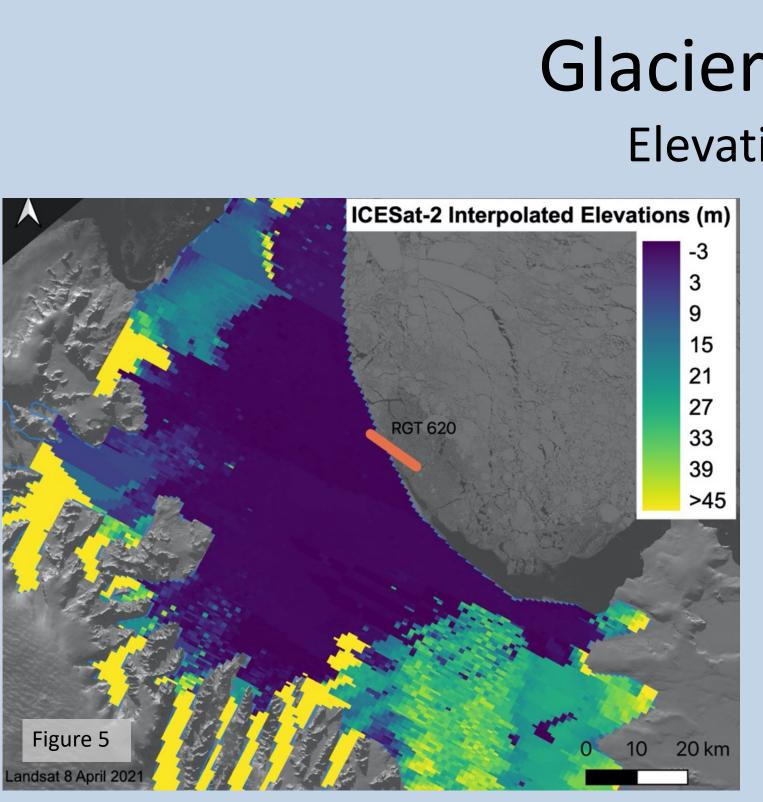




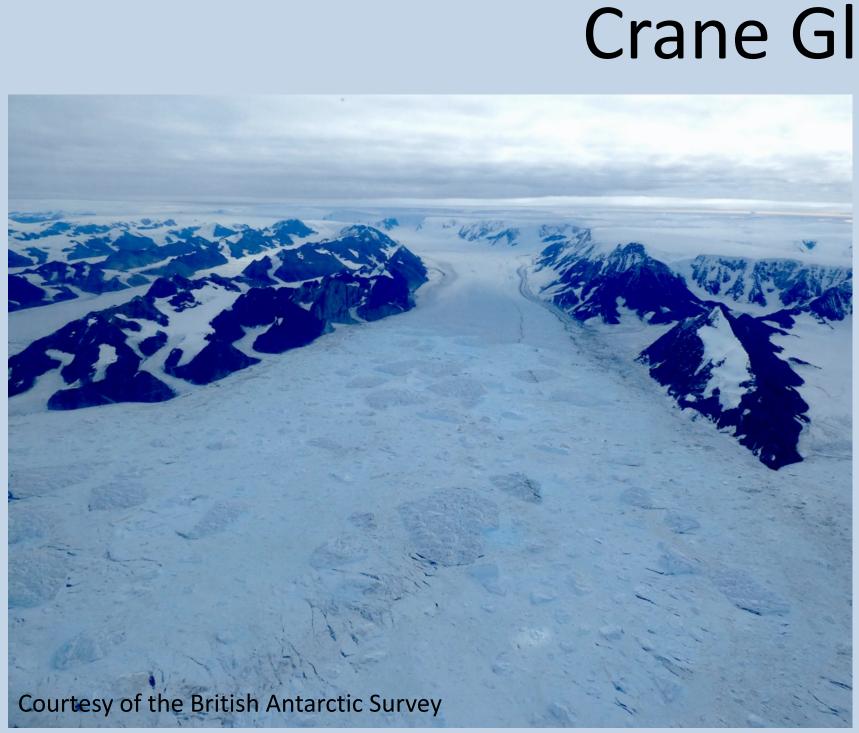
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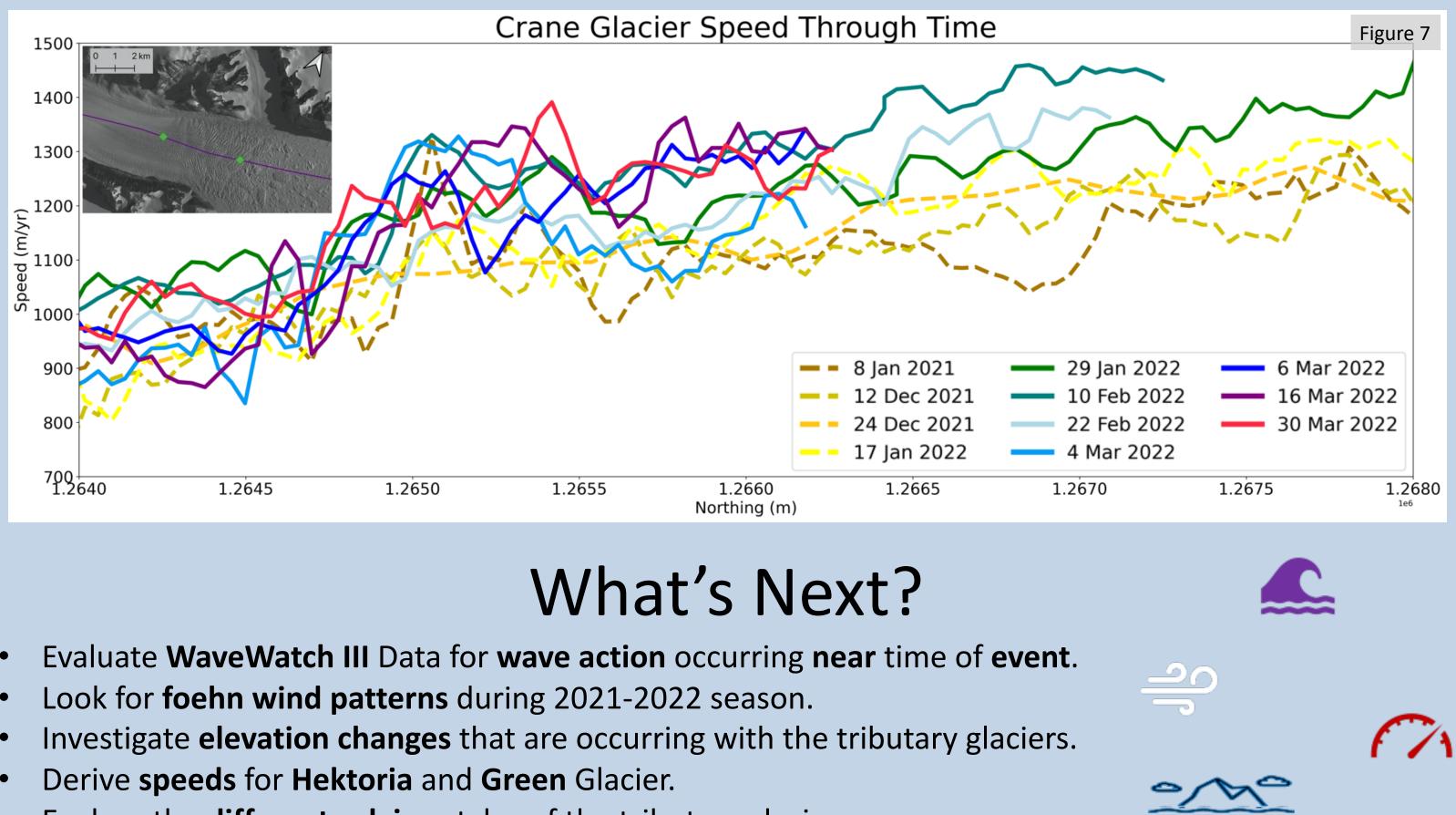


Wave action, (e.g. ocean swells) can weaken the fast ice and catalyze its collapse (Langhorne et al., 2001).



By using an ICESat-2 track that transects open water and landfast ice (orange line in Fig. 5), the minimum thickness of landfast sea ice was calculated to be approximately **6.5 m** (Figure 6).

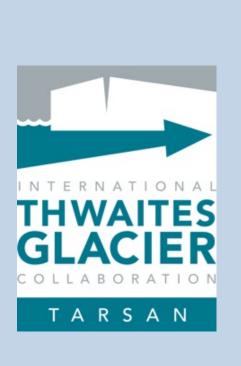




- Explore the **different calving** styles of the tributary glaciers.

ind Marinsek, S.: The role of föhn winds in eastern Antarctic Peninsula rapid ice shelf collapse, The Cryosphere, 16, 1369–1381, 2022. https://doi.org/10.5194/tc-16-1369-2022 Langhorne PJ, Squire VA, Fox C, Haskell TG. Lifetime estimation for a land-fast ice sheet subjected to ocean swell, Annals of Glaciology, 33:333-8, 2001.





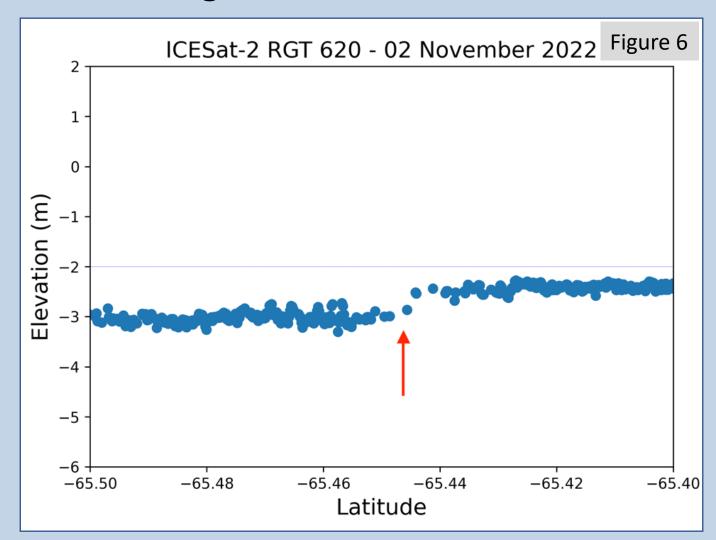


Glacier Response

Elevation Changes

Figure 5: ICESat-2 elevations from 2021 season, prior to landfast ice break-out event.

Initial elevations will be used to estimate the landfast sea ice and tributary glacier thickness, and elevation changes.



Crane Glacier

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Figure 7: Crane Glacier speeds derived from Sentinel 1 speckle tracking using the Alaska Satellite Facility HyP3 pipeline, between the two green dots (upper left).

Dashed lines show speeds from previous months and years prior to the January 2022 event.

Solid lines show Crane Glacier's response to the **removal** of the landfast ice.

Crane's speed has increased on the order of 200-300 m/yr beyond its "normal" seasonal fluctuation.