

HOOCH₂SCHO formation in the marine boundary layer: Rate coefficient measurement of the intramolecular Hydrogen-shift in the CH₃SCH₂OO radical

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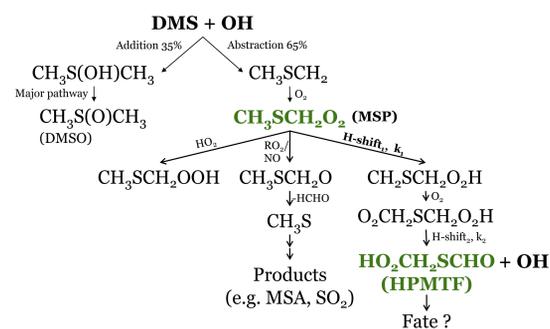
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Introduction

- DMS (CH₃SCH₃) is a major tropospheric source of sulfur and sulfate aerosol in the atmosphere.
- The kinetics and mechanism of DMS oxidation is, therefore, of critical importance to developing a comprehensive understanding of climate and climate change.

Simplified OH initiated DMS atmospheric oxidation mechanism:



What is the fate of CH₃SCH₂O₂ (MSP)?

- high-NO_x: CH₃SCH₂O₂ + NO reaction dominant
- low-NO_x (e.g. remote marine environment): H-shift reaction pathways possible

Previous studies:

- Wu et al. (2014): Theoretical study
 - Proposed H-shift reactions, hydroperoxymethyl thioformate (HOOCH₂SCHO, HPMTF) formation
 - $k_1(298\text{ K}) = 2.1\text{ s}^{-1}$
 - $k_2(298\text{ K}) \gg k_1(298\text{ K})$
- Veres et al. (2019): ATOM field campaign
 - I-ToF-CIMS detection
 - Observed HPMTF in the marine boundary layer
 - $k_1(298\text{ K}) = 0.04\text{ s}^{-1}$ (calculated theoretically)
- Berndt et al. (2019): Experimental study
 - free-jet flow system coupled to I-CI-API-TOF mass spectrometer
 - $k_1(298\text{ K}) = 0.23\text{ s}^{-1}$

Motivations and objectives

The H-shift chemical pathway is currently not included in CCM atmospheric models
 ⇒ potential impact on sulfate aerosol formation and climate

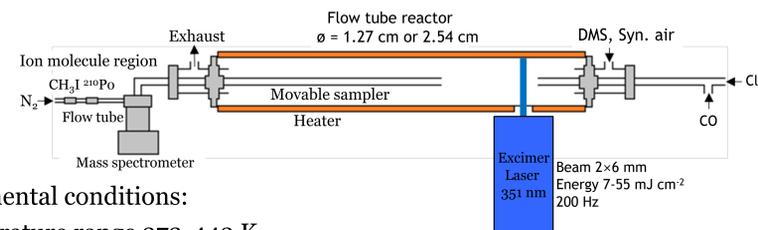
The discrepancy between the calculated and measured rate coefficient calls for additional experimental measurement of k_1 and its temperature dependence

In this study:

- k_1 was measured over a range of temperature using a photolysis flow tube reactor coupled to a chemical ionization mass spectrometer
- $k_1(T)$ was theoretically investigated with the Mo6-2X/aug-cc-pVTZ level of density functional theory and W1 energies

Experimental details

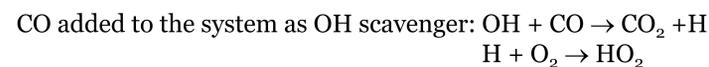
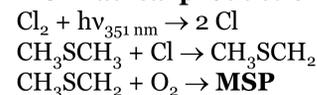
- Pulsed laser photolysis used to generate Cl radicals in presence of excess DMS in a slow-flow tube reactor
- Formation of HPMTF measured with a quadrupole or a high-resolution time-of-flight chemical ionization mass spectrometer (quad-CIMS or HR-ToF-CIMS)



Experimental conditions:

- Temperature range 373-443 K
- Pressure 620 Torr
- [DMS] = 4-12 × 10¹⁴ molecule cm⁻³, [Cl₂] = 2-7 × 10¹¹ molecule cm⁻³, [Cl]₀ = 0.2-4 × 10¹⁰ molecule cm⁻³, [O₂] = 4 × 10¹⁸ molecule cm⁻³, [CO] = 2 × 10¹⁷ molecule cm⁻³

MSP radical production:

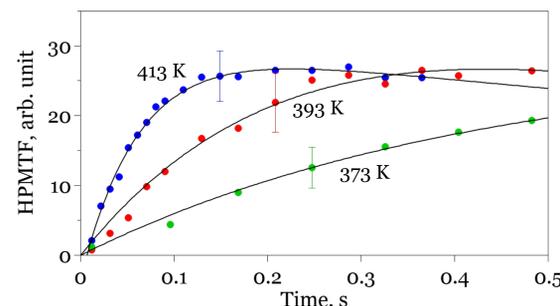


Low radical concentrations:

MSP + MSP and MSP + HO₂ minimized (confirmed with box model calculations)

Results

Representative HPMTF formation profiles



Different conditions tested for each T:

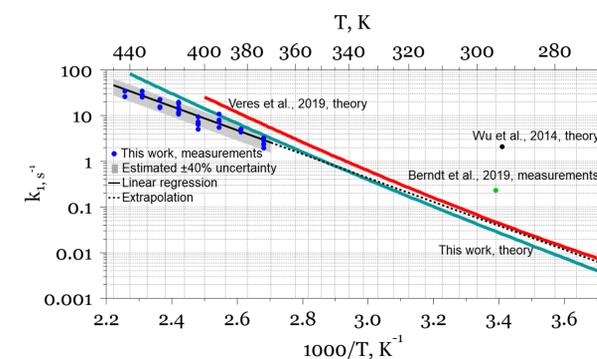
- concentration of reactants
- laser energy
- flow tube reactor diameter
- thermal decomposition of HPMTF
- ⇒ No impact on obtained $k_1(T)$

Lines are fits of the HPMTF profiles for pseudo-first order formation and loss of HPMTF

$$S(t) = S_{\text{asy}} \times \frac{k_1(T)}{k_{\text{loss}} - k_1(T)} \times (\exp(-k_1(T)t) - \exp(-k_{\text{loss}}t))$$

where S_{asy} is the asymptote value of HPMTF and k_{loss} is the rate coefficient of HPMTF first-order loss

Arrhenius plot of $k_1(T)$, CH₃SCH₂O₂ → CH₂SCH₂OOH



Experimental (373-443 K):

$$k_1(T) = 2.98 \times 10^7 \exp\left(-\frac{6020}{T}\right) \text{ s}^{-1}$$

$$k_1(298\text{ K}) = 0.05\text{ s}^{-1} \text{ (extrapolated value)}$$

Theory:

$$k_1(T) = 8120 \times \frac{T}{298} \times 9.62 \exp\left(-\frac{3670}{T}\right) \text{ s}^{-1}$$

$$k_1(298\text{ K}) = 0.037\text{ s}^{-1}$$

- ⇒ agreement between our experiment extrapolation and theory
- agreement with Veres et al. (theory)
- significantly less than Wu et al. (theory) and Berndt et al. (experiments)

Atmospheric Implications

Marine boundary layer loss processes for CH₃SCH₂O₂ (MSP)

Reactant	k # cm ³ molecule ⁻¹ s ⁻¹	Concentration * molecule cm ⁻³	First Order Rate Coefficient s ⁻¹
NO	1 × 10 ⁻¹¹	10 ⁸ -10 ⁹	0.001-0.01
HO ₂	1 × 10 ⁻¹¹	10 ⁸ -10 ⁹	0.001-0.01
RO ₂	1 × 10 ⁻¹¹	10 ⁸ -10 ⁹	0.001-0.01
OH	1 × 10 ⁻¹⁰	10 ⁶	0.0001
H-shift (k_1)	-	-	0.05

Estimated values for NO/HO₂/RO₂ reactions; OH rate coefficient from Fittschen (2019)
 * Estimated values vary with season and location

- H-shift reaction is fast enough to impact the fate of CH₃SCH₂O₂ in remote environments, where NO_x concentrations are low
- Fate of HPMTF needs to be determined to fully evaluate the impact of the H-shift reaction pathway on sulfate aerosol formation and climate

References:

- Wu, R., S. Wang, and L. Wang (2015), New mechanism for the atmospheric oxidation of dimethyl sulfide. The importance of intramolecular hydrogen shift in a CH₃SCH₂OO radical, *J. Phys. Chem. A*, 119(1), 112-117, doi:10.1021/jp511616j
- T. Berndt et al. (2019), Fast Peroxy Radical Isomerization and OH Recycling in the Reaction of OH Radicals with Dimethyl Sulfide, *J. Phys. Chem. Lett.*, 10 (21), 6478-6483, doi:10.1021/acs.jpclett.9b02567
- P. R. Veres et al. (2019), Global aircraft sampling reveals a new dimethyl sulfide oxidation mechanism in the marine atmosphere, submitted to *Proc. Natl. Acad. Sci. U.S.A.*
- C. Fittschen (2019), The reaction of peroxy radicals with OH radicals, *Chem. Phys. Lett.*, 725, 102-108, doi:10.1016/j.cplett.2019.04.002

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